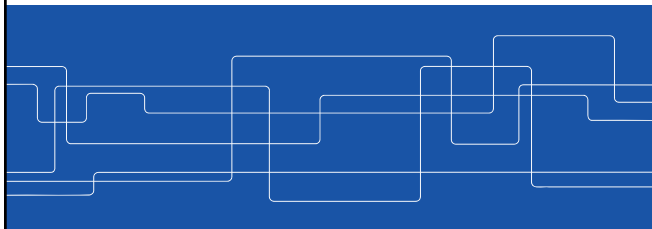




Nuclear Fuel Cycle 2013

Lecture 3: Basic Nuclear Chemistry, Part 2

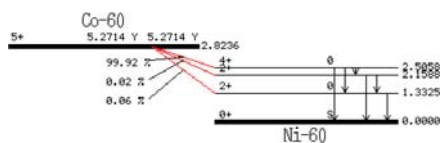




Radioactive decay

$N \rightarrow \text{Daughter} + \text{particle}$ $t_{1/2}$ [s]

- The energy of the mother is higher than that of the daughter.
- The difference in energy is transferred
 - to the particle (as kinetic energy; velocity)
 - and often also to the daughter.
- The daughter loses the "remaining" energy in one or more γ -photons





Radioactive decay

$N \rightarrow \text{Daughter} + \text{particle}$ λ [s^{-1}]


The disappearance of N is a 1st order reaction with respect to N

$$A = -\frac{dN}{dt} = \lambda N \quad \text{The Activity equals the rate of disappearance of N}$$

$$-\int_{N_0}^N \frac{1}{N} dN = \int_0^t \lambda dt \Rightarrow \ln N - \ln N_0 = -\lambda t$$

$$N = N_0 e^{-\lambda t} \quad N = \frac{N_0}{2} \Rightarrow t_{1/2} = \frac{\ln N_0 - \ln\left(\frac{N_0}{2}\right)}{\lambda} = \frac{\ln 2}{\lambda}$$

$$A = A_0 e^{-\lambda t} = A_0 e^{-\frac{\ln 2}{t_{1/2}} t}$$

 **Radioactive equilibrium**

Often the daughter is also radioactive:
 Nuclide1 ($t_{1/2,1}$, λ_1) \rightarrow Nuclide2 ($t_{1/2,2}$, λ_2) \rightarrow Nuclide3


Nuclide1
 \downarrow $t_{1/2,1}$
 Nuclide2
 \downarrow $t_{1/2,2}$
 Nuclide3 (stable)

$$\frac{dN_2}{dt} = -\frac{dN_1}{dt} - \lambda_2 N_2 = \lambda_1 N_1 - \lambda_2 N_2$$

$$N_1 = N_{1(0)} e^{-\lambda_1 t}$$

$$\frac{dN_2}{dt} + \lambda_2 N_2 - \lambda_1 N_{1(0)} e^{-\lambda_1 t} = 0$$

$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) + N_{2(0)} e^{-\lambda_2 t}$$

 **Radioactive equilibrium**

Nuclide1 ($t_{1/2,1}$, λ_1) \rightarrow Nuclide2 ($t_{1/2,2}$, λ_2) \rightarrow Nuclide3


$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) + N_{2(0)} e^{-\lambda_2 t}$$

Assuming that at $t=0$ a quantitative separation between Nuclide 1 and 2 has been achieved, then $N_{2(0)}=0$ and

$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t})$$

Expressed as radioactivity, this is

$$A_2 = \frac{\lambda_2}{\lambda_2 - \lambda_1} A_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) \quad (\text{since } A_i = N_i \lambda_i)$$

 **Radioactive equilibrium ($(t_{1/2})_1 \gg (t_{1/2})_2$)**

$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) + N_{2(0)} e^{-\lambda_2 t}$$

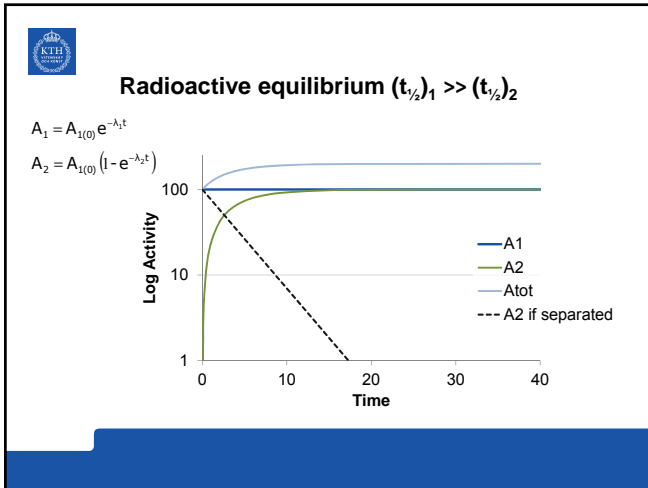
If $(t_{1/2})_1 \gg (t_{1/2})_2$ then $\lambda_2 \gg \lambda_1$ and the equilibrium can be simplified to

$$N_2 = \frac{\lambda_1}{\lambda_2} N_{1(0)} (1 - e^{-\lambda_2 t}) \quad A_2 = A_{1(0)} (1 - e^{-\lambda_2 t})$$

At equilibrium
 $A_1 = A_2 = N_1 \lambda_1 = N_2 \lambda_2 (= N_3 \lambda_3 \dots)$

Example:

^{238}U	$\xrightarrow{\alpha}$	^{234}U	$\xrightarrow{\alpha}$	^{230}Th	$\xrightarrow{\alpha}$	^{226}Ra	$\xrightarrow{4\alpha + 2\beta}$	^{210}Pb	$\xrightarrow{2\beta + \alpha}$	^{206}Pb
$4.5 \times 10^9 \text{y}$		$2.5 \times 10^5 \text{y}$		$7.5 \times 10^4 \text{y}$		$1.6 \times 10^3 \text{y}$		22y		



Radioactive equilibrium $(t_{1/2})_1 > (t_{1/2})_2$

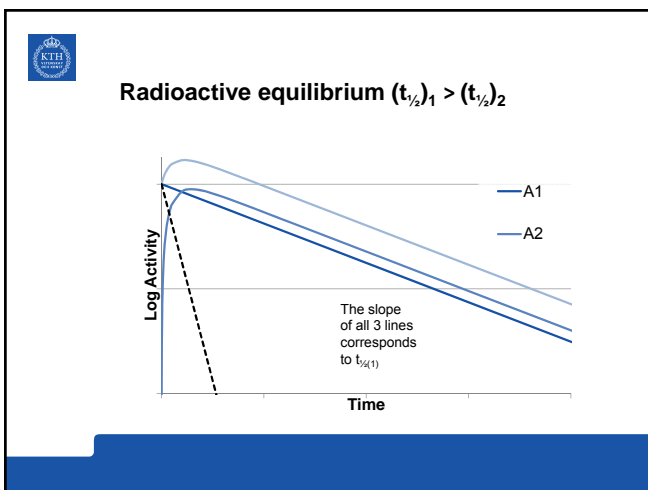
“The whole expression” without simplifications must be used:

$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) \quad \text{or} \quad A_2 = \frac{\lambda_2}{\lambda_2 - \lambda_1} A_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t})$$

Transient equilibrium is reached when

$e^{-(\lambda_2 - \lambda_1)t}$ approaches 0, i.e. when $\lambda_2 t$ can be neglected compared with $\lambda_1 t$

$$\frac{N_2}{N_1} = \frac{\lambda_1}{\lambda_2 - \lambda_1} = \frac{(t_{1/2})_2}{(t_{1/2})_1 - (t_{1/2})_2}$$





Radioactive equilibrium ($t_{1/2,1} < t_{1/2,2}$)

"The whole expression" without simplifications must be used:

$$N_2 = \frac{\lambda_1}{\lambda_2 - \lambda_1} N_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) \quad \text{or} \quad A_2 = \frac{\lambda_2}{\lambda_2 - \lambda_1} A_{1(0)} (e^{-\lambda_1 t} - e^{-\lambda_2 t})$$

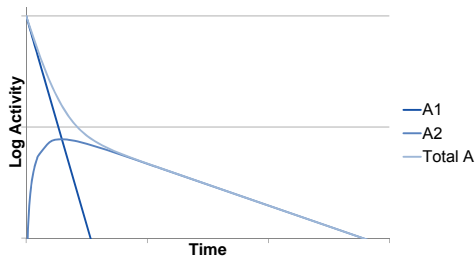
Equilibrium is never reached,

but when $e^{-(\lambda_1 - \lambda_2)t} \ll 1$,

only the decay of the daughter is seen



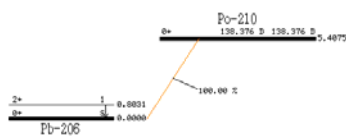
Radioactive equilibrium ($t_{1/2,1} < t_{1/2,2}$)





^{210}Po production

How much pitchblende containing 40% U_3O_8 is needed to produce 1 gram ^{210}Po ? (assuming that ^{210}Po can be extracted quantitatively)



²¹⁰Po production

How much pitchblende containing 40% U₃O₈ is needed to produce 1 gram ²¹⁰Po? (assuming that ²¹⁰Po can be extracted quantitatively)

²¹⁰Po production

Secular equilibrium: $A_1 = A_2 = N_1\lambda_1 = N_2\lambda_2 (= N_3\lambda_3 \dots)$
 $N_{Po-210}\lambda_{Po-210} = N_{U-238}\lambda_{U-238}$

$$N_{Po-210} = \frac{1}{210} \times 6.022 \times 10^{23} = 2.87 \times 10^{21} \text{ atoms}$$

$$t_{1/2, Po-210} = 138.38 \text{ d} \Rightarrow \lambda_{Po-210} = \ln 2 / 138.38 = 5.01 \times 10^{-3} \text{ d}^{-1}$$

$$\lambda_{U-238} = \ln 2 / (4.5 \times 10^9 \times 365) = 4.25 \times 10^{-13} \text{ d}^{-1}$$

$$N_{U-238} = \frac{2.87 \times 10^{21} \times 5.01 \times 10^{-3}}{4.25 \times 10^{-13}} = 3.38 \times 10^{31} \text{ atoms}$$

$$N_{U-238} = \frac{3.38 \times 10^{31}}{6.023 \times 10^{23}} = 5.61 \times 10^7 \text{ mol} \Rightarrow N_{U_3O_8} = \frac{5.61 \times 10^7}{3} = 1.87 \times 10^7 \text{ mol}$$

= 15 750 tonnes U₃O₈ and 39 400 tonnes pitchblende

Example ($t_{1/2,1} < t_{1/2,2}$)

A has a radioactive daughter and a stable grand-daughter
 $A \rightarrow B \rightarrow C(\text{stable})$
 The total activity of this system has been measured. At $t=0$ the activity of B=0. Determine the half-life of B.

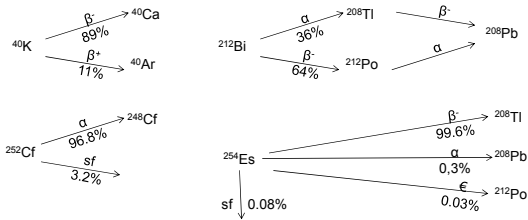
Time [min]	Tot A [Bq]	In A [-]
0	400.0	5.99
1	304.4	5.72
2	235.4	5.46
3	185.4	5.22
4	148.7	5.00
5	121.7	4.80
6	101.6	4.62
7	86.3	4.46
8	74.6	4.31
9	65.5	4.18
10	58.2	4.08
15	36.7	3.60
20	25.4	3.24
25	18.1	2.90
30	13.0	2.56

$t_{1/2} = \ln(2) / 0.066 = 10.5 \text{ min.}$



Dual (triple) decay

Some radionuclides can decay in several ways. Examples:





Dual (triple, etc.) decay

For a dual decay, the decay law is written:



The formation of B and C can be expressed as

$$\frac{dN_B}{dt} = \lambda_B N_A \quad \text{and} \quad \frac{dN_C}{dt} = \lambda_C N_A$$

The decay constant can be written $\lambda = \lambda_1 + \lambda_2 + \dots + \lambda_i = \sum_1^i \lambda_i$

The partial half-lives are related $\frac{1}{(t_{1/2})} = \frac{1}{(t_{1/2})_1} + \frac{1}{(t_{1/2})_2} + \dots + \frac{1}{(t_{1/2})_i} = \sum_1^i \frac{1}{(t_{1/2})_i}$



Radionuclides in nature

Primordial ("original"):
 $t_{1/2} > 10^8$ years

- **Cosmogenic**
 Cosmic radiation: γ -photons, mesons, neutrons, protons, α -particles, heavier particles

- **Anthropogenic**
 Fission products



Primordial radionuclides for Z<82 (Pb)

Nuclide	Isotopic abundance %	Decay mode	$t_{1/2}$, years
⁴⁰ K	0.0117	β^- , EC	1.26×10^9
⁵⁰ V	0.25	β^- , EC	$>1.4 \times 10^{17}$
⁸⁷ Rb	27.83	β^-	4.88×10^{10}
¹¹⁵ In	95.72	β^-	4.4×10^{14}
¹²³ Te	0.905	EC	1.3×10^{13}
¹³⁸ La	0.092	β^- , EC	1.05×10^{11}
¹⁴⁴ Nd	23.8	α	2.1×10^{15}
¹⁴⁷ Sm	15	α	1.06×10^{11}
¹⁴⁸ Sm	11.3	α	7×10^{15}
¹⁷⁶ Lu	2.59	β^-	3.8×10^{10}
¹⁷⁴ Hf	0.162	α	2×10^{15}
¹⁸⁷ Re	62.6	β^-	4.2×10^{10}
¹⁹⁰ Pt	0.012	α	6.5×10^{11}



Cosmogenic radionuclides

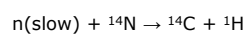
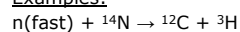
Inflow:

$$R_i = \sigma_i \cdot \Phi_{\text{cosm.}} \cdot N_i$$

/
/

Cross section
Flux

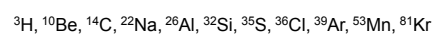
Examples:



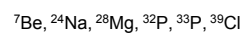


Cosmogenic radionuclides

Long-lived cosmogenic radionuclides appearing in meteorites and rain water



Short-lived cosmogenic radionuclides appearing in rain water





Radiometric dating -Accumulation

Determine age of geologic samples

Decay of mother nuclide: $A = A_0 e^{-\lambda t}$

Build up of daughter: $B = A_0(1 - e^{-\lambda t})$

$$\Rightarrow t = \frac{\ln(1 + B/A)}{\lambda}$$

- t and $t_{1/2}$ should be in same order of magnitude
- Mother and daughter must be stuck in matrix



Radiometric dating

Th-series: ^{232}Th ($t_{1/2} = 1.41 \times 10^{10}\text{y}$) $\xrightarrow{6\alpha, 4\beta^-}$ ^{208}Pb

Np-series: ^{237}Np ($t_{1/2} = 2.1 \times 10^6\text{y}$) $\xrightarrow{7\alpha, 4\beta^-}$ ^{209}Bi

U-series: ^{238}U ($t_{1/2} = 4.5 \times 10^9\text{y}$) $\xrightarrow{8\alpha, 6\beta^-}$ ^{206}Pb

^{235}U ($t_{1/2} = 7.08 \times 10^8\text{y}$) $\xrightarrow{7\alpha, 4\beta^-}$ ^{207}Pb

^{206}Pb , ^{207}Pb and ^{208}Pb are of radiogenic origin while ^{204}Pb is not.

U and Th containing minerals can be dated by measuring the ratio of Pb and mother nuclide.

Another way is to measure the Pb-ratio: $\frac{^{207}\text{Pb}}{^{206}\text{Pb}} = \frac{1}{138} \frac{e^{\lambda_{235}t} - 1}{e^{\lambda_{238}t} - 1}$



Radiometric dating

^{87}Rb ($t_{1/2} = 1.41 \times 10^{10}\text{y}$) \rightarrow ^{87}Sr

^{40}K ($t_{1/2} = 1.27 \times 10^9\text{y}$) \rightarrow ^{40}Ar (most of the argon in the atmosphere has been created this way)

^{187}Re ($t_{1/2} = 5 \times 10^{10}\text{y}$) \rightarrow ^{187}Os

$^{87}\text{Rb}/^{87}\text{Sr}$ is the most reliable method of dating old minerals

The oldest material found in the Earth's crust are 3.5×10^9 years
 Meteoritic stones have been dated 4.5×10^9 years

An approximate age of our solar system (5.9×10^9 years) can be calculated assuming that the ratio $^{235}\text{U}/^{238}\text{U}$ was 1



U-238/U-235 = 1

$$N_{238} = N_{238}^0 e^{-\lambda_{238} t} \implies N_{238}^0 = \frac{N_{238}}{e^{-\lambda_{238} t}}$$

$$N_{238}^0 = N_{235}^0 \implies \frac{N_{238}}{e^{-\lambda_{238} t}} = \frac{N_{235}}{e^{-\lambda_{235} t}}$$

$$\frac{N_{238}}{N_{235}} = \frac{e^{-\lambda_{238} t}}{e^{-\lambda_{235} t}}$$

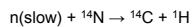
$$\ln\left(\frac{N_{238}}{N_{235}}\right) = (\lambda_{235} - \lambda_{238}) t$$

$$t = \frac{\ln\left(\frac{N_{238}}{N_{235}}\right)}{\lambda_{235} - \lambda_{238}} = 5.9 \times 10^9 \text{ y}$$



Radiocarbon dating

The ^{14}C -method ($t_{1/2}=5736\text{y}$) is by far the most used for determining the age of biologic material.



- It is assumed that the neutron flux has been constant.
=> the production of ^{14}C has been constant
=> equilibrium between production and decay of ^{14}C
=> constant $^{14}\text{CO}_2$ in atmosphere
- When organism dies it ceases to take in new ^{14}C -isotopes



Example

Before 1952 (when atmospheric nuclear bomb tests started) the specific activity of ^{14}C was ≈ 15 decays/min.g (today it's closer to 20 dpm/g).

In a sample containing 250 mg carbon, 2480 C-14 decays occurred during 20 hours. How old is the sample? [$t_{1/2, \text{C-14}} = 5736\text{y}$]

$$\text{Specific activity} = 2480 / (20 \cdot 60 \cdot 0.25) = 8.267 \text{ dpm/g}$$

$$A = A_0 e^{-\lambda t} \implies t = \frac{\ln\left(\frac{A_0}{A}\right)}{\lambda} = \frac{\ln\left(\frac{15}{8.267}\right)}{\ln 2 / 5736} = 4930$$

Fission

Typically 200 MeV is released in a fission event

Fission

Spontaneous fission

- Induced fission
 - When an isotope is struck by a fast neutron: **fissionable**
 - When an isotope is struck by a thermal (slow) neutron: **fissile**

(²³⁵U and ²³⁹Pu are fissionable and fissile)

Nuclear chain reaction

A nuclear chain reaction will take place if sufficient amount of neutrons are produced and captured.

Critical mass: the smallest amount of fissile material needed for a sustained nuclear chain reaction



Calculation example

What is the decay rate of K-40 in 1 g of natural K?

Nuclide	Isotopic abundance %	Decay mode	$t_{1/2}$, years
^{40}K	0.0117	β^- , EC	1.26×10^9

$$A = N\lambda$$

$$N_{^{40}\text{K}} = \frac{1}{M_{\text{K}}} \times I_{^{40}\text{K}} \times 6.023 \times 10^{23} = \frac{1}{39.1} \times 0.000117 \times 6.023 \times 10^{23} = 1.80 \times 10^{18}$$

$$\lambda_{^{40}\text{K}} = \frac{\ln 2}{1.26 \times 10^9 \times 365.25 \times 24 \times 3600}$$

$$A_{^{40}\text{K}} = N_{^{40}\text{K}} \lambda_{^{40}\text{K}} = 31,4 \text{ Bq}$$
